NATIONAL GEOPHYSICAL DATA CENTER
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MONTHLY SUMMARY OF GEOMAGNETIC ACTIVITY

THE GEOMAGNETIC FIELD. The intensity of the Earth's magnetic field at any point in space and time arises from the MAIN field internal to the planet; from the electrical currents flowing in the ionized upper atmosphere; and from the currents induced within the Earth's crust. The main field component varies slowly in time and can be grossly described as a bar magnetic with north and south poles that extend well out into space.

The main field creates a cavity in interplanetary space called the magnetosphere, where the Earth's magnetic field dominates any field carried by the charged particles of the solar wind. The magnetosphere shape resembles a comet—a shape owing to the interaction with the solar wind; it is compressed on the side toward the sun and tail—like on the side away from the sun. The magnetosphere also directs the flow of the particles about the Earth.

Particles flowing in the magnetosphere and ionosphere generate currents, which in turn cause variations in the intensity of the Earth's magnetic field. These EXTERNAL currents generate additional currents in the Earth's upper atmospheric layers, which vary on much shorter time scales than the main field, and they create magnetic changes as large as 10% of the main field.

Certain current systems derive their energy from the regular changing solar radiation throughout the day and year. Other irregular current systems obtain their energy from the interaction of the solar wind with the magnetosphere, from the magnetosphere itself, from the interaction between the magnetosphere and the ionosphere, and from the ionosphere itself. Magnetic activity indices, including those reported below, are designed to describe variations in the Earth's main field caused by these irregular currents.

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See back side for definitions of column headings.

DST FOR OCTOBER 1987 has not been received, so we are printing without it. We will send it out soon as we receive it. Thank you for your patience.

- PLANETARY 3-HOUR RANGE INDEX. K-indices isolate solar particle effects on the Earth's magnetism Κp by classifying into disturbance levels the range of variation of the most unsettled horizontal field component during a 3-hour period. Each activity level relates almost logarithmically to its corresponding amplitude. Three-hour indices discriminate conservatively between true magnetic field perturbations and the regular quiet-day variations produced by ionospheric currents.
- PLANETARY A-INDEX. The A-index ranges from 0 to 400. A indices are daily averages of 'a' indices which convert K-values to a linear scale in field units--a scale that measures equivalent disturbance amplitude of a station at which K=9 has a lower limit of 500 nanotesia (nt). Ap is the daily average of A indices from a global array of observatories.
- DISTURBANCE AMPLITUDE-STORM TIME. Dst tracks variations in the solar-induced electric currents flowing about 5.6 Earth radii above the planet's surface. Each hourly value is the average symmetric disturbance amplitude of the horizontal component recorded at four stations, reduced to equatorial changes. Values are given in nt, and they can be either positive or negative; during a storm they become strongly negative.
- PLANETARY DAILY CHARACTER FIGURE. The Cp-figure is a standardized version of the Ci-figure formerly published and is derived from the indices Kp by converting the daily sum of ap into Cp the range 0.0 (quiet) to 2.5 (highly disturbed).
- JULIAN DAY or day of the year. This number resets to 1 after the end of the year, January 1=1.
- Bart DAY NUMBER OF BARTELS 27-DAY CYCLE. The recurrence of geomagnetic activity every 27 days reflects their solar source. J. Bartels defined a series of 27-day periods to track more easily times of unsettled magnetic conditions. He arbitrarily defined his sequence of 27-day intervals to begin in January 1833.
- $\mbox{Q/D}$ MAGNETICALLY QUIET AND DISTURBED DAYS. The following criteria are used to rank selected days of the month from most (Q1) to the least quiet (Q10) and from most (D1) to least disturbed (D5). The following criteria are used in the ranking: the sum of the 8 Kps, the sum of the squares of the of the 8 Kps, and the greatest Kp.
- The aa indices are three-hourly indices computed from K indices of two antipodal observatories (invariant magnetic latitude 50) and provide a quantitative characterization of the magnetic activity. Half-daily and daily values give an estimation of the activity level very close to that obtained with "am" indices. Values are in nanotesla and correspond to the activity level at an invariant magnetic latitude of 50. The aa indices are computed for:

N = daily values for the northern Hemisphere, S = daily values for the Southern Hemisphere, M1, M2 = half-daily values of aa indices for Greenwich day.

What is an index and why are there so many for terrestrial magnetism? An index continuously summarizes a complex measurement; its discrete values simplify and clarify the variations. Ideally each geomagnetic index should follow a single class of magnetic disturbance; in reality few do. The bewildering array of magnetic indices reflects many past attempts to define measurements that isolate a single source of variation.

The subscript "p" means planetary and designates a global magnetic activity index. The following 13 observatories, which lie between 46 and 63 north and south geomagnetic latitude, now contribute to the planetary indices: Lerwick (UK), Eskdalemuir (UK), Hartland (UK), Ottawa (Canada), Fredericksburg (USA), Meanook (Canada), Sitka (USA), Eyre-well (New Zealand), Canberra (Australia), Lovo (Sweden), Rude Skov (Denmark), Wingst (Germany), and Witteveen (the Netherlands).

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